

VISIBLE AND NEAR-IR REFLECTANCE SPECTRA FOR SMECTITE, SULFATE AND PERCHLORATE UNDER DRY CONDITIONS FOR INTERPRETATION OF MARTIAN SURFACE MINERALOGY. R. V. Morris¹, D. C. Golden², D. W. Ming¹, T. G. Graff², R. E. Arvidson³, S. M. Wiseman³, K. A. Lichtenberg³, and S. Cull³. ¹ARES NASA Johnson Space Center, Houston TX 77058 (richard.v.morris@nasa.gov), ²ESCG-JW23, Houston, TX, ³Washington University, St. Louis, MO.

Introduction: Visible and near-IR (VNIR) spectral data for the martian surface obtained from orbit by the MRO-CRISM and OMEGA instruments are interpreted as having spectral signatures of H₂O/OH-bearing phases, including smectites and other phyllosilicates, sulfates, and high-SiO₂ phases [e.g., 1-4]. Interpretations of martian spectral signatures are based on and constrained by spectra that are obtained in the laboratory on samples with known mineralogical compositions and other physicochemical characteristics under, as appropriate, Mars-like environmental conditions (e.g., temperature, pressure, and humidity). With respect to environmental conditions, differences in the absolute concentration of atmospheric H₂O can effect the hydration state and therefore the spectra signatures of smectite phyllosilicates (solvation H₂O) and certain sulfates (hydration H₂O) [e.g., 5-7].

We report VNIR spectral data acquired under humid (laboratory air) and dry (dry N₂ gas) environments for two natural smectites (nontronite API-33A and saponite SapCa-1) to characterize the effect of solvation H₂O on spectral properties. We also report spectral data for the thermal dehydration products of (1) melanterite (FeSO₄·7H₂O) in both air and dry N₂ gas and (2) Mg-perchlorate (Mg(ClO₄)₂·6H₂O) in dry N₂ environments. Spectral measurements for samples dehydrated in dry N₂ were made without exposing them to humid laboratory air.

Methods: VNIR reflectivity spectra (0.35-2.50 μ m) were acquired with Analytical Spectral Devices (ASD) fiberoptic spectrometers configured with Mug lights. A FieldSpec3 Hi-Res spectrometer was used for measurements in laboratory air at ~25°C. A FieldSpec3 spectrometer, a hot plate (ambient to 400°C), and a dewpoint meter (Vaisala DRYCAP DM70) were co-located in a one-atmosphere glove box that can be kept under continuous purge using dry N₂ gas (from liquid N₂) to provide a dry atmosphere. Spectral measurements on samples in the glove box were made in air (8000-12000 ppm H₂O by volume) prior to starting the N₂ purge, after purging 24-650 hr at ~25°C (150-400 ppm H₂O or 0.5-2.0% relative humidity), and, depending on the sample, after heating to 50-330°C on the hot plate and cooling to ~25°C. Because H₂O evolves during heating, the time at temperature was governed by the time required for the H₂O content of the gas at-

mosphere to return to <400 ppm.

Powder X-ray diffraction spectra (PANalytical X'Pert PRO) were obtained under ambient conditions for all samples within 5-10 min and ~1 hr after their removal from the glove box or after thermal treatment in air to determine mineralogical composition and phase stability.

Results and Discussion:

Nontronite and Saponite. Dehydration of nontronite and saponite under hyper-arid conditions results in a dramatic decrease in intensity of spectral features that require the presence of only the H₂O molecule (e.g., at ~1.9 μ m) relative to those that require the presence of only M-OH (e.g., at 2.29 and 2.31 μ m for nontronite (M = Fe) and saponite (M = Mg), respectively (Fig. 1)). All M-OH spectral features between 2.0 and 2.5 μ m are also better resolved with increasing loss of solvation H₂O. Spectral features that are composites of H₂O-only and M-OH-only absorptions will also decrease in intensity and can shift in position if the individual peaks occur at different wavelengths. Saponite is an example, where a broad band centered near 1.41 μ m (Fig. 1b, blue) under humid conditions is sharp and centered at 1.39 μ m for measurements under hyper-arid conditions (red).

The enhancement of M-OH relative to H₂O spectral features in Fig. 1 is in agreement with previous studies of smectites [e.g., 6-8]. The magnitude of the enhancements cannot be directly compared, however, because environmental H₂O concentrations are not reported in all the earlier work.

If climatic conditions on Mars favor loss of solvation H₂O over geologic time, these results imply that spectral features for M-OH should dominate over those for H₂O for smectites and other phases that can exchange H₂O with the martian atmosphere. We suggest that dehydration of solvation H₂O from smectites by heating at ~110°C for laboratory timescales is a reasonable surrogate for dehydration of solvation H₂O over geologic timescales at lower temperatures.

Fe-Sulfates. The first thermal decomposition product of melanterite in both air and dry N₂ atmospheres is szomolnokite (FeSO₄·H₂O) (Fig. 2). The spectra are similar to literature szomolnokite spectra [e.g., 9], except that no spectral feature at ~0.43 μ m from Fe³⁺ is present. In humid air at higher temperatures (~240°C),

Fe^{2+} oxidation occurs resulting in formation of a ferric sulfate hydroxide which is possibly hydrated (from XRD data). The spectral features from $\text{Fe}^{3+}\text{-OH}$ are located at 1.49, 1.82, 2.23, and 2.37 μm . This phase may occur naturally on Mars [10]. Under hyper arid (and presumably low O_2) conditions, szomolnokite thermally altered to crystalline FeSO_4 (from XRD) (Fig. 2b). The broad bands centered near 0.99 and 1.53 μm are from Fe^{2+} electronic transitions.

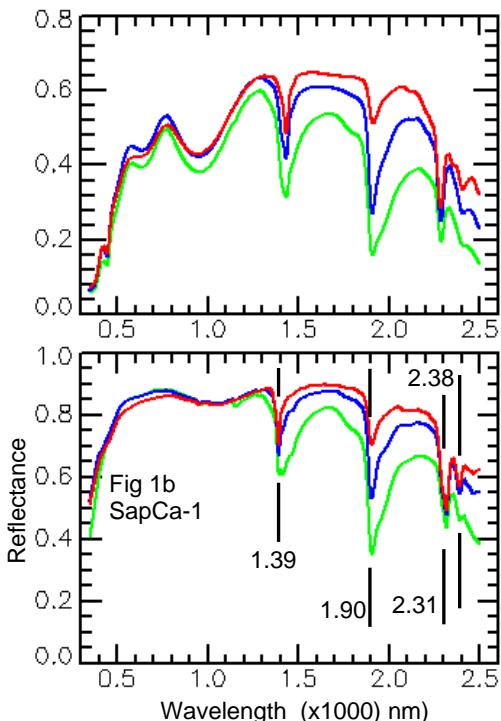


Fig. 1. (a) Nontronite (API-33A) spectra in air (green), after 24 hr N₂ purge at 25°C with final H₂O = 266 ppm (blue), and after 24 hr N₂ purge at 110°C with final H₂O = 200 ppm (red). (b) Saponite SapCa-1 spectra at ambient conditions (green), after 24 hr N₂ purge at 25°C with final H₂O = 266 ppm (blue), and after 24 hr N₂ purge at 110°C with final H₂O = 200 ppm (red).

Mg-Perchlorate. The three most intense spectral features for fully and partially hydrated Mg-perchlorate are located at 1.42, 1.91, and 2.38 μm (Fig. 3). Because these positions approximate those for hydrous ferric sulfate, it is possible that locations previously interpreted as polyhydrated sulfates are actually mixed (and hydrated sulfate and perchlorate salts).

References: [1] Bibring *et al.* (2006). *Science*, 312, 400. [2] Poulet *et al.* (2005), *Nature*, 438, 623. [3] Mustard *et al.*, *Nature*, doi:10.1038/nature07097, 305. [4] Milliken *et al.* (2008) *Geology*, 36, 847. [5] Bishop *et al.* (1994) *Clays Clay*

Min., 42, 702. [6] Milliken and Mustard (2005) *JGR*, 110, E12001, doi:10.1029/2005JE002534. [7] Chipera and Vaniman (2007) *GCA*, 71, 241. [8] Xu *et al.* (2000) *Clays Clay Min.*, 48, 120. [9] Crowley *et al.* (2003) *GEEA*, 3, 219. [10] Lichtenberg *et al.*, this volume.

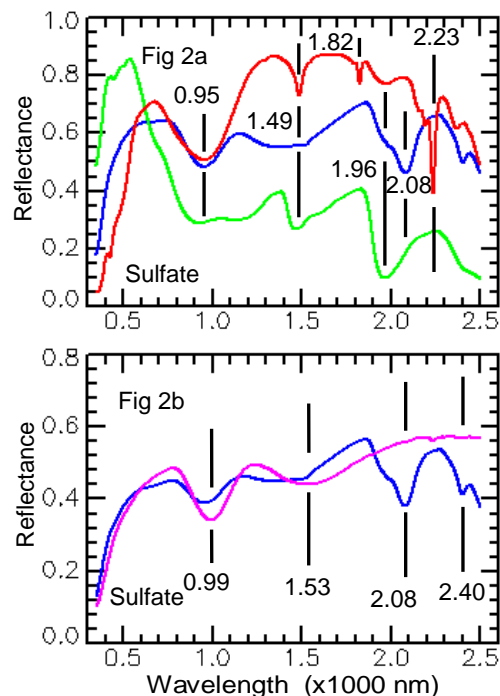


Fig. 2. (a) Spectra of melanterite ($\text{FeSO}_4 \cdot 7\text{H}_2\text{O}$) (green) and szomolnokite ($\text{FeSO}_4 \cdot \text{H}_2\text{O}$) (blue) and ferric sulfate hydroxide (red) obtained by thermal dehydration in air. (b) Spectra of szomolnokite (blue) and anhydrous FeSO_4 (magenta) obtained as thermal dehydration products of melanterite under dry conditions at 210°C and 310°C, respectively.

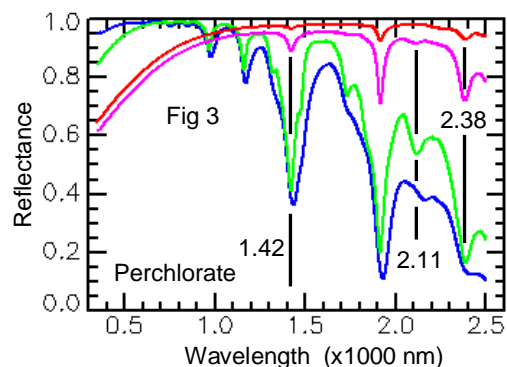


Fig 3. Spectra of unheated $\text{Mg}(\text{ClO}_4)_2 \cdot 6\text{H}_2\text{O}$ (blue) and its thermal dehydration products in dry N₂ at 210°C (green) and 330°C (red and magenta).